## Time-Stepping Methods for PDEs and Ocean Models

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Scientific Computing

## Outline

1. On the Spatial and Temporal Order of Convergence of PDEs
1.1 Analytical Derivation
1.2 Derivation by Symbolic Algebra
1.3 Numerical Experiments
2. Time-Stepping Methods for Ocean Models
2.1 Barotropic-Baroclinic Splitting and Filtering of Barotropic Modes
2.2 Verification Suite of Shallow Water Test Cases

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## 1. On the Spatial and Temporal Order of Convergence of PDEs

Pop Quiz: Order of Convergence of Global Solution Error Norm with Respect to Exact Solution

$$
\text { You are modeling the PDE } u_{t}=\mathcal{F}\left(u, u_{x}, u_{x x}, \cdots, x, t\right)
$$

| Numerical Method | Refinement in Space: | Refinement in Time: | Refinement in Space and Time: |
| :--- | :--- | :--- | :--- |
| $\mathcal{O}\left(\Delta x^{\alpha}\right), \mathcal{O}\left(\Delta t^{\beta}\right)$ | $\Delta x \rightarrow 0, \Delta t$ fixed | $\Delta t \rightarrow 0, \Delta x$ fixed | $\Delta x \rightarrow 0, \Delta t \rightarrow 0, \Delta t / \Delta x$ fixed |
| $\alpha=2, \beta=1$ |  |  |  |
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| $\alpha=2, \beta=3$ |  |  |  |

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- With a stable numerical scheme, the order of accuracy of the global solution error is the same as that of the global truncation error,

$$
\hat{\tau}_{G}=\mathcal{O}\left(\Delta x^{\alpha}\right)+\Delta t \mathcal{O}\left(\Delta x^{\alpha}\right)+\Delta t^{2} \mathcal{O}\left(\Delta x^{\alpha}\right)+\cdots+\Delta t^{\beta-1} \mathcal{O}\left(\Delta x^{\alpha}\right)+\mathcal{O}\left(\Delta t^{\beta}\right)
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which can be approximated as

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\hat{\tau}_{G} \approx \mathcal{O}\left(\Delta x^{\alpha}\right)+\mathcal{O}\left(\Delta t^{\beta}\right), \text { for } \Delta t \ll 1
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- A simultaneous refinement of $\Delta t$ and $\Delta x$, while maintaining their ratio $\Delta t / \Delta x=\gamma$, a constant, yields

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$$

- Strategy: Given $\alpha$, we need $\beta \geq \alpha$ to obtain maximum possible order of accuracy. But we gain no improvement in order of convergence for $\beta>\alpha$ despite more work. So, optimum choice is $\beta=\alpha$.


## 1. On the Spatial and Temporal Order of Convergence of PDEs

Order of convergence of the error norm in the asymptotic regime at constant ratio of time-step to grid spacing for varying orders of spatial and temporal discretizations

| Order of Spatial Discretization $\alpha$ | Time-Stepping Method Employed | Order of Time-Stepping Method $\beta$ | Order of Convergence of Error Norm in Asymptotic Regime at Constant Ratio of Time-Step to Grid Spacing $\min (\alpha, \beta)$ |
| :---: | :---: | :---: | :---: |
| 1 | FE | 1 | $\min (1,1)=1$ |
| 1 | RK2 or AB2 | 2 | $\min (1,2)=1$ |
| 1 | RK3 or AB3 | 3 | $\min (1,3)=1$ |
| 1 | RK4 or AB4 | 4 | $\min (1,4)=1$ |
| 2 | FE | 1 | $\min (2,1)=1$ |
| 2 | RK2 or AB2 | 2 | $\min (2,2)=2$ |
| 2 | RK3 or AB3 | 3 | $\min (2,3)=2$ |
| 2 | RK4 or AB4 | 4 | $\min (2,4)=2$ |
| 3 | FE | 1 | $\min (3,1)=1$ |
| 3 | RK2 or AB2 | 2 | $\min (3,2)=2$ |
| 3 | RK3 or AB3 | 3 | $\min (3,3)=3$ |
| 3 | RK4 or AB4 | 4 | $\min (3,4)=3$ |
| 4 | FE | 1 | $\min (4,1)=1$ |
| 4 | RK2 or AB2 | 2 | $\min (4,2)=2$ |
| 4 | RK3 or AB3 | 3 | $\min (4,3)=3$ |
| 4 | RK4 or AB4 | 4 | $\min (4,4)=4$ |

FE $\equiv$ forward Euler, RK $\equiv$ Runge-Kutta, and $A B \equiv$ Adams-Bashforth

## 1. On the Spatial and Temporal Order of Convergence of PDEs: Motivation

- A graduate level textbook on numerical analysis typically contains standard predictor-corrector and multistep time-stepping methods applied to ODEs in one chapter, followed by spatial discretization operators of PDEs in another.
- In real-world applications, the discretization of the PDE consists of both spatial and temporal components.
- The order of convergence of a PDE with spatial and/or temporal refinement is a function of both the mesh spacing $\Delta x$ and the time step $\Delta t$.
- I investigate this simultaneous dependence of the local truncation error of the numerical solution of a PDE on $\Delta x$ and $\Delta t$, for varying orders of spatial and temporal discretizations.


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### 1.1. Analytical Derivation of Local Truncation Error

## Local Truncation Error of a Generic Hyperbolic PDE

Theorem 1. Given the exact solution $u_{j}^{n}$ of a hyperbolic PDE $u_{t}=\mathcal{F}\left(u, u_{x}, x, t\right)$ on a uniform mesh with spacing $\Delta x$, at spatial locations $x_{j}$ for $j=1,2, \ldots$, and at time level $t^{n}$, the exact solution at time level $t^{n+1}=t^{n}+\Delta t$ may be obtained by Taylor expanding $u_{j}^{n}$ about time level $t^{n}$ as

$$
u_{j}^{n+1}=u_{j}^{n}+\sum_{k=1}^{\infty} \frac{\Delta t^{k}}{k!}\left(\frac{\partial^{k} u}{\partial t^{k}}\right)_{j}^{n} \equiv u_{j}^{n}+\sum_{k=1}^{\infty} \frac{\Delta t^{k}}{k!}\left(\mathcal{F}^{(k)}\right)_{j}^{n},
$$

where $\left(\mathcal{F}^{(k)}\right)_{j}^{n}=\left(\frac{\partial^{k} u}{\partial t^{k}}\right)_{j}^{n}$ is the $k^{\text {th }}$-order spatial derivative at $x_{j}$ and $t^{n}$. The numerical solution at time level $t^{n+1}$, obtained with a time-stepping method belonging to the Method of Lines, may be written in the general form

$$
\hat{u}_{j}^{n+1}=u_{j}^{n}+\sum_{k=1}^{\infty} \frac{\Delta t^{k}}{k!}\left(\widehat{\mathcal{F}}^{(k)}+\mathcal{O}\left(\Delta x^{\alpha}\right)\right)_{j}^{n},
$$

where $\alpha$ is the order of the spatial discretization and $\widehat{\mathcal{F}}^{(k)}$ is specified by the time-stepping method. If $\beta$ represents the order of the time-stepping method,

$$
\left(\widehat{\mathcal{F}}^{(k)}\right)_{j}^{n}=\left(\mathcal{F}^{(k)}\right)_{j}^{n} \equiv\left(\frac{\partial^{k} u}{\partial t^{k}}\right)_{j}^{n}, \text { for } k=1,2, \ldots, \beta
$$

The local truncation error is then

$$
\begin{aligned}
\hat{\tau}_{j}^{n+1} & =u_{j}^{n+1}-\hat{u}_{j}^{n+1} \\
& =\frac{\Delta t}{1!} \mathcal{O}\left(\Delta x^{\alpha}\right)+\frac{\Delta t^{2}}{2!} \mathcal{O}\left(\Delta x^{\alpha}\right)+\frac{\Delta t^{3}}{3!} \mathcal{O}\left(\Delta x^{\alpha}\right)+\cdots+\frac{\Delta t^{\beta}}{\beta!} \mathcal{O}\left(\Delta x^{\alpha}\right)+\frac{\Delta t^{\beta+1}}{(\beta+1)!}\left(c_{\beta+1}+\mathcal{O}\left(\Delta x^{\alpha}\right)\right)_{j}^{n}+\mathcal{O}\left(\Delta t^{\beta+2}\right) \\
& =\Delta t \mathcal{O}\left(\Delta x^{\alpha}\right)+\Delta t^{2} \mathcal{O}\left(\Delta x^{\alpha}\right)+\Delta t^{3} \mathcal{O}\left(\Delta x^{\alpha}\right)+\cdots+\Delta t^{\beta} \mathcal{O}\left(\Delta x^{\alpha}\right)+\mathcal{O}\left(\Delta t^{\beta+1}\right),
\end{aligned}
$$

where $\left(c_{\beta+1}\right)_{j}^{n}=\left(\mathcal{F}^{(\beta+1)}\right)_{j}^{n}-\left(\hat{\mathcal{F}}^{(\beta+1)}\right)_{j}^{n} \neq 0$.
Bishnu, S., Petersen, M., Quaife, B., "On the Spatial and Temporal Order of Convergence of Hyperbolic PDEs", Journal of Computational Physics (submitted)

### 1.1. Analytical Derivation of Local Truncation Error

But wait! I can still verify the order of accuracy by refining only $\Delta x$ or $\Delta t$ !

- Assume a stable numerical scheme, $\Delta t \ll 1$, and the the global solution error is of the same order of accuracy as the global truncation error $\hat{\tau}_{G} \approx \mathcal{O}\left(\Delta x^{\alpha}\right)+\mathcal{O}\left(\Delta t^{\beta}\right) \approx \zeta \Delta x^{\alpha}+\zeta_{\beta+1} \Delta t^{\beta}$.
- Convergent behavior as $\Delta t \rightarrow 0$, keeping $\Delta x$ fixed (refinement only in time)
- Given $\Delta x$ and $\Delta t$, measure the global solution error at a time horizon.
- Reduce $\Delta t$ by a constant ratio, say $p$, but keep $\Delta x$ fixed.
- Measure the global solution error at the same time horizon.
- Plot the norm of the difference between the errors against $\Delta t$.
- Proof: For two time steps $\Delta t_{i}$ and $\Delta t_{i+1}$, with $\Delta t_{i+1} / \Delta t_{i}=p<1$, we can write

$$
\begin{gathered}
\left(\hat{\tau}_{G_{i}}\right)_{j} \approx \zeta \Delta x^{\alpha}+\zeta_{\beta+1} \Delta t_{i}^{\beta}, \quad\left(\hat{\tau}_{G_{i+1}}\right)_{j} \approx \zeta \Delta x^{\alpha}+\zeta_{\beta+1} \Delta t_{i+1}^{\beta} \\
\Delta\left\{\left(\hat{\tau}_{G_{i, i+1}}\right)_{j}\right\}=\left(\hat{\tau}_{G_{i}}\right)_{j}-\left(\hat{\tau}_{G_{i+1}}\right)_{j}=\zeta_{\beta+1}\left(\Delta t_{i}^{\beta}-\Delta t_{i+1}^{\beta}\right)=\zeta_{\beta+1} \Delta t_{i+1}^{\beta}\left(p^{-\beta}-1\right) .
\end{gathered}
$$

Taking logarithm of both sides,

$$
\log \left[\Delta\left\{\left(\hat{\tau}_{G_{i, i+1}}\right)_{j}\right\}\right]=\theta+\beta \log \left(\Delta t_{i+1}\right), \quad \text { where } \theta=\log \left\{\zeta_{\beta+1}\left(p^{-\beta}-1\right)\right\} \text { is constant. }
$$

- Note that the exact solution is independent of $\Delta x$ or $\Delta t$. So,

$$
\Delta \hat{\tau}_{G} \equiv \hat{\tau}_{G^{1}}-\hat{\tau}_{G^{2}}=\left(u_{\text {exact }}-u_{\text {numerical }}^{1}\right)-\left(u_{\text {exact }}-u_{\text {numerical }}^{2}\right)=u_{\text {numerical }}^{2}-u_{\text {numerical }}^{1} .
$$

- By plotting norm of error (or numerical solution) difference between successive spatial resolutions, we can attain convergence with spatial order of accuracy.


### 1.1. Analytical Derivation of Local Truncation Error

## Increase in Global Solution Error with only Temporal Refinement

For certain PDEs and discretization methods, the global solution error can increase with only temporal refinement. A simple example is the one-dimensional linear homogeneous constant-coefficient advection equation $u_{t}+a u_{x}=0$, discretized in space with the first-order upwind finite difference scheme and advanced in time with the first-order Forward Euler method. The global truncation error, approximating the global solution error, is

$$
\left[\left(\hat{\tau}_{G}\right)_{j}\right]_{\text {leading order }}=-\frac{1}{2}|a| \Delta x\left(1-\frac{|a| \Delta t}{\Delta x}\right)\left(u_{x x}\right)_{j}^{n}=-\frac{1}{2}|a| \Delta x(1-C)\left(u_{x x}\right)_{j}^{n}
$$

where $C=|a| \Delta t / \Delta x$ is the Courant number, which is positive and must be less than one to ensure numerical stability. Maintaining $C<1$, if $\Delta x$ is held constant and $\Delta t$ is refined, then $(1-C)$ increases towards 1 , and the magnitude of the global truncation error increases. Moreover, the error will be diffusive in nature.

Numerical Example: $\Delta x=1 / 2^{8}$ (fixed)



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### 1.2. Derivation by Symbolic Algebra

Developed a Symbolic Python (SymPy) library (consisting of $\sim 12,600$ lines of code) that contains

- Taylor Series expansion in $x, y, z$,
- routines for determining the local truncation error of
- the generic ODE $u_{t}=\mathcal{F}(u, t)$, and the generic hyperbolic PDE $u_{t}=\mathcal{F}\left(u, u_{x}, x, t\right)$
- a specific ODE $u_{t}+\left(p_{0}+q_{1}\right) u=f(t)$, and specific PDEs, such as the inhomogeneous, linear variable-coefficient and non-linear advection equations

$$
\begin{aligned}
u_{t}+p(x) u+(q(x) u)_{x} & =f(x, t) \\
u_{t}+u u_{x} & =f(x, t)
\end{aligned}
$$

If $p(x)=p_{0}, q(x)=q_{0}+q_{1} x, u$ and $f$ are only functions of $t$, the linear PDE reduces to the ODE, and so does its truncation errors. I have used

- first-, second-, and third-order spatial discretizations for the PDEs
- five explicit time-stepping methods
- first-order Forward Euler method
- second-order explicit midpoint method
- Williamson's low-storage third-order Runge-Kutta method
- second-order Adams-Bashforth method
- third-order Adams-Bashforth method
- three implicit time-stepping methods
- first-order Backward Euler method
- second-order implicit midpoint method
- second-order Crank-Nicholson method (Trapezoidal Rule)


### 1.2. Derivation by Symbolic Algebra

Relevant Terms in the Local Truncation Error of the Generic One-Dimensional Advection Equation

|  | $\frac{1}{3!} \widehat{\mathcal{F}}^{(3)}$ | $\frac{1}{4} \mathcal{F} \mathcal{F}_{u} \mathcal{F}_{u v} \mathcal{V}+\frac{1}{4} \mathcal{F} \mathcal{F}_{u v} \mathcal{F}_{v} w_{1}+\frac{1}{4} \mathcal{F} \mathcal{F}_{u v} \mathcal{F}_{x}+\frac{1}{8} \mathcal{F}_{t t}$ |
| :---: | :---: | :---: |
| Explicit <br> Midpoint <br> Method | $\frac{1}{3!} c_{3}$ | $\begin{aligned} & \frac{1}{6} \mathcal{F} \mathcal{F}_{u}^{2}+\frac{1}{12} \mathcal{F} \mathcal{F}_{u} \mathcal{F}_{u v} v+\frac{1}{4} \mathcal{F} \mathcal{F}_{u v} \mathcal{F}_{v} w_{1}+\frac{1}{12} \mathcal{F} \mathcal{F}_{u v} \mathcal{F}_{x}+\frac{1}{6} \mathcal{F} \mathcal{F}_{u x} \mathcal{F}_{v}+\frac{1}{6} \mathcal{F}_{t} \mathcal{F}_{u}+\frac{1}{24} \mathcal{F}_{t t} \\ & +\frac{1}{3} \mathcal{F}_{u}^{2} \mathcal{F}_{v} v+\frac{1}{6} \mathcal{F}_{u} \mathcal{F}_{u v} \mathcal{F}_{v} v^{2}+\frac{1}{2} \mathcal{F}_{u} \mathcal{F}_{v}^{2} w_{1}+\frac{1}{6} \mathcal{F}_{u} \mathcal{F}_{v} \mathcal{F}_{v x} v+\frac{1}{3} \mathcal{F}_{u} \mathcal{F}_{v} \mathcal{F}_{x}+\frac{1}{2} \mathcal{F}_{u v} \mathcal{F}_{v}^{2} v w_{1} \\ & +\frac{1}{6} \mathcal{F}_{u v} \mathcal{F}_{v} \mathcal{F}_{x} v+\frac{1}{3} \mathcal{F}_{u x} \mathcal{F}_{v}^{2} v+\frac{1}{6} \mathcal{F}_{v}^{3} w_{2}+\frac{1}{2} \mathcal{F}_{v}^{2} \mathcal{F}_{v x} w_{1}+\frac{1}{6} \mathcal{F}_{v}^{2} \mathcal{F}_{x x}+\frac{1}{6} \mathcal{F}_{v} \mathcal{F}_{v x} \mathcal{F}_{x}+\frac{1}{6} \mathcal{F}_{v} \mathcal{F}_{x t} \end{aligned}$ |
|  | $\frac{1}{3!} \widehat{\mathcal{F}}^{(3)}$ | $\begin{aligned} & \frac{1}{4} \mathcal{F} \mathcal{F}_{u}^{2}+\frac{1}{4} \mathcal{F} \mathcal{F}_{u} \mathcal{F}_{u v} v+\frac{1}{2} \mathcal{F} \mathcal{F}_{u v} \mathcal{F}_{v} w_{1}+\frac{1}{4} \mathcal{F} \mathcal{F}_{u v} \mathcal{F}_{x}+\frac{1}{4} \mathcal{F} \mathcal{F}_{u x} \mathcal{F}_{v}+\frac{1}{4} \mathcal{F}_{t} \mathcal{F}_{u}+\frac{1}{8} \mathcal{F}_{t t} \\ & +\frac{1}{2} \mathcal{F}_{u}^{2} \mathcal{F}_{v} v+\frac{1}{4} \mathcal{F}_{u} \mathcal{F}_{u v} \mathcal{F}_{v} v^{2}+\frac{3}{4} \mathcal{F}_{u} \mathcal{F}_{v}^{2} w_{1}+\frac{1}{4} \mathcal{F}_{u} \mathcal{F}_{v} \mathcal{F}_{v x} v+\frac{1}{2} \mathcal{F}_{u} \mathcal{F}_{v} \mathcal{F}_{x}+\frac{3}{4} \mathcal{F}_{u v} \mathcal{F}_{v}^{2} v w_{1} \\ & +\frac{1}{4} \mathcal{F}_{u v} \mathcal{F}_{v} \mathcal{F}_{x} v+\frac{1}{2} \mathcal{F}_{u x} \mathcal{F}_{v}^{2} v+\frac{1}{4} \mathcal{F}_{v}^{3} w_{2}+\frac{3}{4} \mathcal{F}_{v}^{2} \mathcal{F}_{v x} w_{1}+\frac{1}{4} \mathcal{F}_{v}^{2} \mathcal{F}_{x x}+\frac{1}{4} \mathcal{F}_{v} \mathcal{F}_{v x} \mathcal{F}_{x}+\frac{1}{4} \mathcal{F}_{v} \mathcal{F}_{x t} \end{aligned}$ |
| Midpoint <br> Method | $\frac{1}{3!} c_{3}$ | $\begin{aligned} & -\frac{1}{12} \mathcal{F} \mathcal{F}_{u}^{2}+\frac{1}{12} \mathcal{F} \mathcal{F}_{u} \mathcal{F}_{u v} v+\frac{1}{12} \mathcal{F} \mathcal{F}_{u v} \mathcal{F}_{x}-\frac{1}{12} \mathcal{F} \mathcal{F}_{u x} \mathcal{F}_{v}-\frac{1}{12} \mathcal{F}_{t} \mathcal{F}_{u}+\frac{1}{24} \mathcal{F}_{t t}-\frac{1}{6} \mathcal{F}_{u}^{2} \mathcal{F}_{v} v \\ & -\frac{1}{12} \mathcal{F}_{u} \mathcal{F}_{u v} \mathcal{F}_{v} v^{2}-\frac{1}{4} \mathcal{F}_{u} \mathcal{F}_{v}^{2} w_{1}-\frac{1}{12} \mathcal{F}_{u} \mathcal{F}_{v} \mathcal{F}_{v x} v-\frac{1}{6} \mathcal{F}_{u} \mathcal{F}_{v} \mathcal{F}_{x}-\frac{1}{4} \mathcal{F}_{u v} \mathcal{F}_{v}^{2} v w_{1}-\frac{1}{12} \mathcal{F}_{u v} \mathcal{F}_{v} \mathcal{F}_{x} v \\ & -\frac{1}{6} \mathcal{F}_{u x} \mathcal{F}_{v}^{2} v-\frac{1}{12} \mathcal{F}_{v}^{3} w_{2}-\frac{1}{4} \mathcal{F}_{v}^{2} \mathcal{F}_{v x} w_{1}-\frac{1}{12} \mathcal{F}_{v}^{2} \mathcal{F}_{x x}-\frac{1}{12} \mathcal{F}_{v} \mathcal{F}_{v x} \mathcal{F}_{x}-\frac{1}{12} \mathcal{F}_{v} \mathcal{F}_{x} t \end{aligned}$ |

Recall that for second-order time-sepping methods, $\widehat{\mathcal{F}}^{(1)}=\mathcal{F}^{(1)}, \widehat{\mathcal{F}}^{(2)}=\mathcal{F}^{(2)}$, but $\widehat{\mathcal{F}}^{(3)} \neq \mathcal{F}^{(3)}$ leading to $c_{3}=\mathcal{F}^{(3)}-\widehat{\mathcal{F}}^{(3)} \neq 0$.

### 1.2. Derivation by Symbolic Algebra

Terms containing $\Delta t^{\prime} \Delta x^{k} I \times k \in\{\{1,2\} \times\{0,1,2\}\} \cup\{\{3\} \times\{0\}\}$ within local truncation error of the numerical solution of the linear inhomogeneous variable-coefficient advection equation $u_{t}+p(x) u+(q(x) u)_{x}=f(x, t)$, discretized in space with first order upwind finite difference and advanced in time with explicit midpoint method

| $l$ | $k$ | Term containing $\Delta t^{t} \Delta x^{k}$ within the Local Truncation Error |
| :---: | :---: | :---: |
| 1 | 0 | 0 |
|  | 1 | $\Delta t\left[\Delta x\left\{-\frac{1}{2} q u_{x x}-q_{x} u_{x}-\frac{1}{2} q_{x x} u+\ldots\right\}\right]$ |
|  | 2 | $\Delta t\left[\Delta x^{2}\left\{\frac{1}{6} q u_{x x x}+\frac{1}{2} q_{x} u_{x x}+\frac{1}{2} q_{x x} u_{x}+\frac{1}{6} q_{x x x} u+\ldots\right\}\right]$ |
| 2 | 0 | 0 |
|  | 1 | $\begin{aligned} \Delta t^{2}[\Delta x\{ & -\frac{1}{4} f q_{x x}-\frac{1}{2} f_{x} q_{x}-\frac{1}{4} f_{x x} q+\frac{1}{2} p q u_{x x}+p q_{x} u_{x}+\frac{1}{2} p q_{x x} u \\ & +\frac{1}{2} p_{x} q u_{x}+\frac{1}{2} p_{x} q_{x} u+\frac{1}{4} p_{x x} q u+\frac{1}{2} q^{2} u_{x x x}+\frac{9}{4} q q_{x} u_{x x} \\ & \left.\left.+\frac{7}{4} 4 q_{x x} u_{x}+\frac{1}{2} q q_{x x} u+\frac{3}{2} q_{x}^{2} u_{x}+q_{x} q_{x x} u+\ldots\right\}\right] \end{aligned}$ |
|  | 2 | $\begin{aligned} & \Delta t^{2}\left[\Delta x ^ { 2 } \left\{\frac{1}{12} f q_{x x x}+\frac{1}{4} f_{x} q_{x x}+\frac{1}{4} f_{x x} q_{x}+\frac{1}{12} f_{x x} q-\frac{1}{6} p q u_{x x x}-\frac{1}{2} p q_{x} u_{x x}-\frac{1}{2} p q_{x x} u_{x}\right.\right. \\ & \quad-\frac{1}{6} p q_{x x x} u-\frac{1}{4} p_{x} q u_{x x}-\frac{1}{2} p_{x} q_{x} u_{x}-\frac{1}{4} p_{x} q_{x x} u-\frac{1}{4} p_{x x} q u_{x}-\frac{1}{4} p_{x x} q_{x} u-\frac{1}{12} p_{x x x} q u \\ &\left.\left.\quad-\frac{7}{4} q q_{x} u_{x x x}-\frac{17}{8} q q_{x x} u_{x x}-\frac{5}{4} q q_{x x x} u_{x}-\frac{7}{4} q_{x}^{2} u_{x x}-\frac{5}{2} q_{x} q_{x x} u_{x}-\frac{2}{3} q_{x} q_{x x x} u-\frac{3}{8} q_{x x}^{2} u+\ldots\right)\right] \end{aligned}$ |
| 3 | 0 | $\begin{aligned} \Delta t^{3} & {\left[\frac{1}{6} f p^{2}+\frac{1}{3} f p q_{x}+\frac{1}{6} f p_{x} q+\frac{1}{6} f q q_{x x}+\frac{1}{6} f q_{x}^{2}-\frac{1}{6} f_{t} p-\frac{1}{6} f_{t} q_{x}+\frac{1}{24} f_{t u}+\frac{1}{3} f_{x} p q\right.} \\ & +\frac{1}{2} f_{x} q q_{x}-\frac{1}{6} f_{x} q+\frac{1}{6} f_{x x} q^{2}-\frac{1}{6} p^{3} u-\frac{1}{2} p^{2} q u_{x}-\frac{1}{2} p^{2} q_{x} u-\frac{1}{2} p p_{x} q u-\frac{1}{2} p q^{2} u_{x x} \\ & -\frac{3}{2} p q q_{x} u_{x}-\frac{1}{2} p q q_{x x} u-\frac{1}{2} p q_{x}^{2} u-\frac{1}{2} p_{x} q^{2} u_{x}-\frac{2}{3} p_{x} q q_{x} u-\frac{1}{6} p_{x x} q^{2} u-\frac{1}{6} q^{3} u_{x x x} \\ & \left.-q^{2} q_{x} u_{x x}-\frac{2}{3} q^{2} q_{x x} u_{x}-\frac{1}{6} q^{2} q_{x x x} u-\frac{7}{6} q q_{x}^{2} u_{x}-\frac{2}{3} q q_{x} q_{x x} u-\frac{1}{6} q_{x}^{3} u+\ldots\right] \end{aligned}$ |


By specifying all spatial gradients to zero, the local truncation error reduces to that of the ODE $u_{t}+\left(p_{0}+q_{1}\right) u=f(t)$, advanced with the explicit midpoint method, $\Delta t^{3}\left[\frac{1}{6} f p_{0}^{2}+\frac{1}{3} f p_{0} q_{1}+\frac{1}{6} f q_{1}^{2}-\frac{1}{6} f_{t} p_{0}-\frac{1}{6} f_{t} q_{1}+\frac{1}{24} f_{\text {tt }}-\frac{1}{6} p_{0}^{3} u-\frac{1}{2} p_{0}^{2} q_{1} u-\frac{1}{2} p_{0} q_{1}^{2} u-\frac{1}{6} q_{1}^{3} u\right]+\mathcal{O}\left(\Delta t^{4}\right)$

## Outline

1. On the Spatial and Temporal Order of Convergence of PDEs
1.1 Analytical Derivation
1.2 Derivation by Symbolic Algebra
1.3 Numerical Experiments
2. Time-Stepping Methods for Ocean Models
2.1 Barotropic-Baroclinic Splitting and Filtering of Barotropic Modes
2.2 Verification Suite of Shallow Water Test Cases

### 1.3. Numerical Experiments: Linear Advection

Convergence of Linear Advection using First-Order Upwind (Finite Difference) in Space ( $\alpha=1$ )


Convergence of Linear Advection using Piecewise Parabolic Reconstruction (Finite Volume) in Space ( $\alpha \approx 3$ )


### 1.3. Numerical Experiments: Non-Linear Burgers' Advection

Convergence of Non-Linear Advection using First-Order Upwind (Finite Difference) in Space ( $\alpha=1$ )


Convergence of Non-Linear Advection using Piecewise Parabolic Reconstruction (Finite Volume) in Space ( $\alpha \approx 3$ )


Number of cells


Number of cells


Number of time steps

## Outline

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### 2.1. Barotropic-Baroclinic Splitting

- Ocean circulation models deals with disparate time scales by splitting the momentum equations into two parts:
- a barotropic part for solving the depth independent fast 2D barotropic waves (advanced in time either explicitly using a small time-step or implicitly using a long time-step) and
- a baroclinic part for solving the much slower 3D baroclinic waves
- Before reconciling the barotropic variables with their baroclinic counterparts to arrive at the total 3D states, a time-averaging filter is applied over the barotropic solutions, to minimize aliasing and mode-splitting errors.


Barotropic-Baroclinic Splitting of Velocity: $u=\bar{u}+u^{\prime}$


2D Barotropic Part $\bar{u}$
3D Baroclinic Part $u^{\prime}$ with Vertical Mean $\vec{u}=0$

Kang, H., Evans, K., Petersen, M., Jones, P., and Bishnu, S., (2021), "A scalable semi-implicit barotropic mode solver for the MPAS-Ocean", Journal of Advances in Modeling Earth Systems

### 2.1. Time-Averaging Filters Incorporated in MPAS-Ocean

Rectangular and Cosine Filters with Primary (Red) \& Secondary (Blue) Weights


Rectangular Filter of Range 0.5


Cosine Filter


Hamming Window and Shchepetkin's Filters with Primary (Red) \& Secondary (Blue) Weights


Second Order Accurate Filter of Shchepetkin et al.


Minimal Dispersion Filter of Shchepetkin et al.

### 2.1. Surface Gravity Wave Simulation in MPAS-Ocean with Various Filters

Numerical SSH with RK4 vs split-explicit method using rectangular and cosine filters


Numerical SSH with RK4 vs split-explicit method using Hamming Window and Shchepetkin's filters


RK4 vs Split-Explicit Method with Second-Order Accurate Filter of Shchepetkin et al.


RK4 vs Split-Explicit Method with Minimal Dispersion Filter of Shchepetkin et al.


### 2.1. Shallow Water Solver Simulating Surface Gravity Wave

- To understand the combined stabilizing effect of various barotropic time-averaging filters and the forward-backward (FB) parameters, I developed a non-linear shallow water solver in object-oriented Python and tested it against the simulation of a surface gravity wave.
- I obtain a near-exact solution using a truncated Fourier series approximation, which is spectrally accurate in space, and the classic fourth-order Runge-Kutta (RK4) method in time. I treat it as the reference benchmark to compare to my numerical solution, employing piecewise parabolic reconstruction in space and the forward-backward (FB) time-stepping method with parameter $\gamma$,
$u^{n+1}=u^{n}+\mathcal{F}\left(u^{n}, \eta^{n}\right) \Delta t ; \eta^{n+1}=\eta^{n}+\left\{(1-\gamma) \mathcal{G}\left(u^{n}, \eta^{n}\right)+\gamma \mathcal{G}\left(u^{n+1}, \eta^{n}\right)\right\} \Delta t$, where $u_{t}=\mathcal{F}(u, \eta) ; \eta_{t}=\mathcal{G}(\eta, t)$ represent the non-linear shallow water equations in functional form.
- The following table lists maximum error norms of the surface elevation of the gravity wave after 1 hour ( 30 baroclinic time steps, each consisting of 2 minutes and 20 barotropic subcycles) for a variety of filters and FB parameter $\gamma$.


## Surface Elevation Maximum Error Norm $\times 10^{-3}$

| FB | No | Rectangular Filter with Range $R$ |  |  |  | Cosine Filters |  |  | Shchepetkin Filters |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Parameter $\gamma$ | Filter | $R=0.25$ | $R=0.375$ | $R=0.50$ | $R=0.75$ | $R=1.00$ | ROMS | HW | $2^{\text {nd }}$ Order | Min. Disp. |
| -0.50 | 2.322 | 2.611 | 1.780 | 2.277 | $\mathbf{2 . 4 0 8}$ | $\mathbf{3 . 0 3 9}$ | 1.945 | 1.573 | $\mathbf{1 . 9 5 1}$ |  |
| -0.25 | 2.192 | 2.514 | 1.703 | 2.191 | 2.441 | 3.073 | 1.834 | 1.452 | 1.976 |  |
| +0.00 | 2.065 | 2.417 | 1.607 | 2.107 | 2.474 | 3.121 | 1.737 | 1.333 | 2.003 | 1.685 |
| +0.25 | 1.946 | 2.327 | 1.521 | $\mathbf{2 . 0 7 5}$ | 2.506 | 3.238 | 1.641 | 1.230 | 2.035 | 1.741 |
| +0.50 | 1.847 | 2.240 | 1.453 | 2.101 | 2.537 | 3.354 | 1.554 | 1.138 | 2.088 |  |
| +0.75 | 1.750 | 2.154 | $\mathbf{1 . 3 8 6}$ | 2.126 | 2.567 | 3.470 | 1.486 | 1.048 | 2.141 | 1.902 |
| +1.00 | 1.653 | 2.070 | 1.461 | 2.151 | 2.601 | 3.592 | $\mathbf{1 . 4 1 8}$ | $\mathbf{1 . 0 1 6}$ | 2.197 | 2.083 |
| +1.25 | 1.583 | 1.986 | 1.542 | 2.184 | 2.640 | 3.718 | 1.477 | 1.106 | 2.274 | 2.182 |
| +1.50 | $\mathbf{1 . 5 1 5}$ | $\mathbf{1 . 9 6 4}$ | 1.628 | 2.220 | 2.678 | 3.843 | 1.558 | 1.214 | 2.352 | 2.284 |

[^0]
## Outline

## 1. On the Spatial and Temporal Order of Convergence of PDEs <br> 1.1 Analytical Derivation <br> 1.2 Derivation by Symbolic Algebra <br> 1.3 Numerical Experiments

2. Time-Stepping Methods for Ocean Models

### 2.1 Barotropic-Baroclinic Splitting and Filtering of Barotropic Modes

2.2 Verification Suite of Shallow Water Test Cases

### 2.2. Verification Suite of Barotropic Test Cases

Motivation: The development of any numerical ocean model warrants a suite of verification exercises for testing its spatial and temporal discretizations. I have designed a set of shallow water test cases for verifying the barotropic solver of ocean models.

## Geophysical Waves and Barotropic Tide

(1) Non-Dispersive Coastal Kelvin Wave
(2) Low Frequency Dispersive Planetary Rossby Wave
(3) Low Frequency Dispersive Topographic Rossby Wave
(9) High Frequency Dispersive Inertia Gravity Wave
(3) Non-Dispersive Equatorial Kelvin Wave

- Dispersive Equatorial Yanai Wave
(1) Low Frequency Dispersive Equatorial Rossby Wave
(8) High Frequency Dispersive Equatorial Inertia Gravity Wave
(0) Barotropic Tide


## Standard Mathematical Test Cases

(1) Diffusion Equation
(2) Viscous Burgers Equation
(3) Non-linear Manufactured Solution

### 2.2. Verification Suite of Barotropic Test Cases

I developed a new unstructured-mesh ocean model (consisting of $\sim 12,600$ lines of code) in object-oriented Python, employing TRiSK-based spatial discretization, and the following set of time-stepping algorithms:

## Standard Mathematical Time-Stepping Algorithms

(1) Forward Backward Method or Implicit Euler Method
(2) Explicit Midpoint Method, a Form of Second-Order Runge-Kutta Method
(3) Low-Storage Third-Order Runge-Kutta Method of Williamson
(9) Low-Storage Fourth-Order Runge-Kutta Method of Carpenter and Kennedy
(0) Second-Order Adams-Bashforth Method
( Third-Order Adams-Bashforth Method
(0) Fourth-Order Adams-Bashforth Method

## Time-Stepping Algorithms Popular in Ocean Modeling

(1) Leapfrog Trapezoidal Method
(2) Leapfrog Adams Moulton Method
(3) Forward Backward Method with RK2 Feedback
(9) Generalized Forward Backward Method with AB2 - AM3 Step

- Generalized Forward Backward Method with AB3 - AM4 Step


### 2.2. Verification Suite: Coastal Kelvin Wave




### 2.2. Verification Suite: High-Frequency Inertia-Gravity Wave




### 2.2. Verification Suite: Low-Frequency Planetary Rossby Wave




### 2.2. Verification Suite: Low-Frequency Topographic Rossby Wave




### 2.2. Verification Suite: Barotropic Tide




### 2.2. Verification Suite: Non-Linear Manufactured Solution




### 2.2. Verification Suite: Summary of Shallow Water Test Cases

Summary of Shallow Water Test Cases for the Barotropic Solver of Ocean Models

|  | Coriolis <br> Parameter | Bottom Topography | Numerical PDE | Boundary Conditions |
| :---: | :---: | :---: | :---: | :---: |
| Coastal Kelvin Wave | Constant <br> (f-plane) | Flat Bottom | Linear, Homogeneous, Constant-Coefficient | Non-Periodic in $x$, Periodic in $y$ |
| Inertia-Gravity Wave | Constant <br> (f-plane) | Flat <br> Bottom | Linear, Homogeneous, Constant-Coefficient | Periodic in $x$, <br> Periodic in $y$ |
| Planetary Rossby Wave | Linear in $y$ (beta plane) | Flat Bottom | Linear, Inhomogeneous, Variable-Coefficient | Periodic in $x$, Non-Periodic in $y$ |
| Topographic Rossby Wave | Constant <br> (f-plane) | Linear in $y$, Sloping Bottom | Linear, Inhomogeneous, Variable-Coefficient | Periodic in $x$, Non-Periodic in $y$ |
| Barotropic Tide | Constant (f-plane) | Flat <br> Bottom | Linear, Homogeneous, Constant-Coefficient | Non-Periodic in $x$, Non-Periodic in $y$ |
| Manufactured Solution | Constant <br> (f-plane) | Flat Bottom | Non-Linear, Inhomogeneous, Constant-Coefficient | Periodic in $x$, <br> Periodic in $y$ |

### 2.2. Verification Suite: Convergence of Spatial Operators

## Convergence of TRiSK-based gradient, divergence, curl, and flux interpolation operators




Number of cells

Tangential Velocity along Edges


Number of cells


Number of cells

Curl Operator Interpolated to Cell Centers


Number of cells

## Recap Slide 1. On the Order of Convergence of PDEs

Order of convergence of the error norm in the asymptotic regime at constant ratio of time-step to grid spacing for varying orders of spatial and temporal discretizations

| Order of Spatial Discretization $\alpha$ | Time-Stepping Method Employed | Order of Time-Stepping Method $\beta$ | Order of Convergence of Error Norm in Asymptotic Regime at Constant Ratio of Time-Step to Grid Spacing $\min (\alpha, \beta)$ |
| :---: | :---: | :---: | :---: |
| 1 | FE | 1 | $\min (1,1)=1$ |
| 1 | RK2 or AB2 | 2 | $\min (1,2)=1$ |
| 1 | RK3 or AB3 | 3 | $\min (1,3)=1$ |
| 1 | RK4 or AB4 | 4 | $\min (1,4)=1$ |
| 2 | FE | 1 | $\min (2,1)=1$ |
| 2 | RK2 or AB2 | 2 | $\min (2,2)=2$ |
| 2 | RK3 or AB3 | 3 | $\min (2,3)=2$ |
| 2 | RK4 or AB4 | 4 | $\min (2,4)=2$ |
| 3 | FE | 1 | $\min (3,1)=1$ |
| 3 | RK2 or AB2 | 2 | $\min (3,2)=2$ |
| 3 | RK3 or AB3 | 3 | $\min (3,3)=3$ |
| 3 | RK4 or AB4 | 4 | $\min (3,4)=3$ |
| 4 | FE | 1 | $\min (4,1)=1$ |
| 4 | RK2 or AB2 | 2 | $\min (4,2)=2$ |
| 4 | RK3 or AB3 | 3 | $\min (4,3)=3$ |
| 4 | RK4 or AB4 | 4 | $\min (4,4)=4$ |

FE $\equiv$ forward Euler, RK $\equiv$ Runge-Kutta, and $A B \equiv$ Adams-Bashforth

### 2.2. Verification Suite: Convergence of Shallow Water Test Cases

Convergence of the coastal Kelvin wave, the high-frequency inertia-gravity wave, the barotropic tide, and the non-linear manufactured solution with simultaneous refinement in space and time


## Conclusions, Future Work and Current Status

## Conclusions

## On the Spatial and Temporal Order of Convergence of Hyperbolic PDEs

- The order of convergence at constant ratio of time step to cell width is determined by the minimum of the orders of the spatial and temporal discretizations.
- Convergence of the error norm cannot be guaranteed under only spatial or temporal refinement.


## Time-Stepping Methods for Ocean Models

- The amount of dissipation applied to stabilize the barotropic modes can be controlled by (a) the time-averaging filter, or (b) the forward-backward time-stepping parameters. Too much dissipation can damp the entire solution, not just the spurious oscillations.
- The order of convergence of an ocean model under simultaneous refinement in space and time is limited by minimum of the orders of accuracy of the time-stepping method, and all spatial operators like gradient, divergence, curl etc.


## Ongoing and Future Work

- Extend truncation error analysis and the convergence studies to parabolic equations, higher order and spectral discretizations in space and time, and time integrators beyond Method of Lines.
- Design verification exercises with complexity in between the barotropic and the full primitive equations, involving stratification, a complex bathymetry, and the ability to test both the barotropic and baroclinic components separately.


## Relevant Publications

- Bishnu, S., Petersen, M., Quaife, B., "On the Spatial and Temporal Order of Convergence of Hyperbolic PDEs", Journal of Computational Physics (submitted)
- Bishnu, S., Petersen, M., Quaife, B., "A Suite of Verification Exercises for the Barotropic Solver of Ocean Models" (in preparation)


## Current Status

- Successfully defended PhD Dissertation on June 10, 2021.
- Hoping to continue working at the Los Alamos National Laboratory (LANL) as a postdoctoral researcher and collaborate with scientists working on E3SM at LANL and other national laboratories.


[^0]:    HW $\equiv$ Hamming Window Cosine Filter and Min. Disp $\equiv$ Shchepetkin Filter Optimized for Minimal Numerical Dispersion

